

OSMIUM ISOTOPE CONSTRAINTS ON THE PROPORTION OF BOLIDE COMPONENT IN CHICXULUB IMPACT BRECCIAS

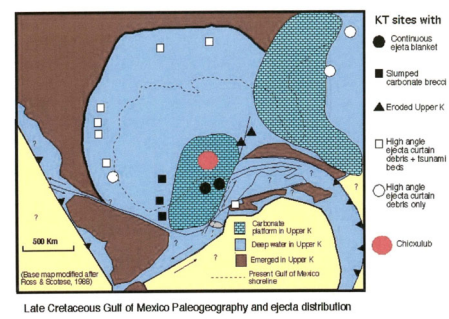


Figure 1: Paleogeography of the Late Cretaceous, [16].

Samples:

Three types of impact breccias have been studied. The first type is a green altered impact meltrock found in the lower portion of the impact sequence. The texture of the rock is microcrystalline and is composed of pyroxene, plagioclase, and alkali feldspars. Its composition is consistent with continental margin rocks. It is generally massive with some flow structures. The rock was brecciated and altered after solidification and contains small amounts of both shocked and unshocked clasts of the impact lithologies. These lithologies include lithic quartzite, and isolated feldspar crystals. The compositions of these rocks are similar to those seen in meltrocks sampled by the Yucatan-6 borehole [8, 9]. Our study includes samples YAX-1 861.4, YAX-1 863.51, YAX-1 876.46, which represent both the top and lower portion of the of the green impact crater. The middle sample in the sequence has the least amount of (mineralogical) alteration [10].

Depth, m	Unit	Sample ID
790	Tertiary Rock	
800	Reworked Suevite	YAX - 1 800.43
810	Suevite	YAX - 1 813.41
820	Suevite	YAX - 1 819.83
830	Brown Impact Melt Breccia	YAX - 1 829.56 YAX - 1 832.83
840	Suevite	YAX - 1 836.34 YAX - 1 841.32
850	Suevite Impact Breccia	YAX - 1 857.65
860	Green Monomict Impact Melt	YAX - 1 861.40 YAX - 1 863.51
870	Green Monomict Impact Melt	YAX - 1 876.46
880	Green Monomict Impact Melt	YAX - 1 883.13
890	Polymict Impact Breccia	
900	Upper K Rock	

The second type of melt breccia under study is a brown altered impact meltrock. It also has a microcrystalline texture and both shocked and unshocked clasts of the target material. Even though this rock type has been altered, remnant schleiren, metagranite, and micritic calcite have been identified. Sample YAX-1 841.32 is representative of this type of rock. It was recovered from a polymictic breccia in the middle of the impact sequence.

Three samples of xenoliths have also been included in the analysis. They also were obtained from the YAX-1 borehole. The first is the sample YAX-1 821.76. It is a mafic clast containing very fine-grained diorite or basaltic andesite. It contains a minor amount of alteration (~ 5%) of the chlorite after mafic sites. The second is a granitoid clast labeled YAX-1 828.28. It is a coarse-grained holocrystalline porphyritic granite. The last xenolith is also a granitoid clast of medium- to coarse-grained holocrystalline equigranular granodiorite. It is significantly altered with carbonate veins, K-feldspar matrix flooding and shredly biotite after black biotite.

Additional samples have been obtained via the Universidad Nacional Autonoma de Mexico (UNAM). These samples contain some of the highest recorded indium concentrations as analyzed by UNAM. It is suspected that they include clasts of the impactor as well as large clasts of feldspars

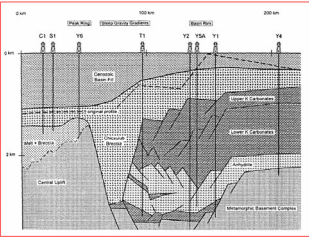


Figure 3: Schematic diagram of Chicxulub impact structure.

The samples come from the borehole marked Y1. The UNAM samples also come from Y1. They are green monomictic impact melt breccias.

Analytical Methods

Samples were spiked for isotope dilution analysis and digested in Carius tubes in reverse *aqua regia* at 230°C. Subsequent separation/ purification was accomplished via solvent extraction (Os) and anion exchange chromatography (Re). Osmium analysis was accomplished via negative thermal ionization mass spectrometry using the University of Maryland Bobcat I mass spectrometer. Rhenium was analyzed using the University of Maryland *Nu Plasma* multi-collector ICP-MS. Analysis was done using a triple electron multiplier arrangement for simultaneous measurement of masses 185, 187, and 190 (for ¹⁸⁷Os correction). Blanks for Re and Os averaged 13 pg, and 1.0 pg, respectively. The ¹⁸⁷Os/¹⁸⁸Os of the blank averaged 0.17. All ratio measurements were better than ± 0.5%. Most samples have quite low concentrations. Analytical uncertainties are variable, with the greatest uncertainty resulting from high blank/sample ratios. Our initial two measurements were of sample aliquants of approximately 0.1 g (YAX-1 861.4; YAX-1 876.46). Due to low concentrations of Re and Os, uncertainties in Os concentrations and isotopic composition are approximately ± 10%. Rhenium quantities were at blank levels and are not reported. More material was digested for subsequent samples and blank/sample ratios decreased accordingly. Uncertainties for these samples are approximately ± 1 for isotopic composition and ± 1-6% and ± 5-30% for Os and Re concentrations, respectively.

The ¹⁸⁷Re/¹⁸⁸Os ratios for all samples are relatively low, so corrections for 65 Ma of ¹⁸⁷Os ingrowth are minor.

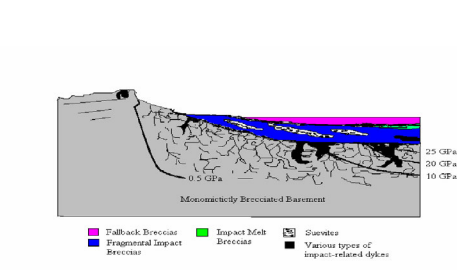
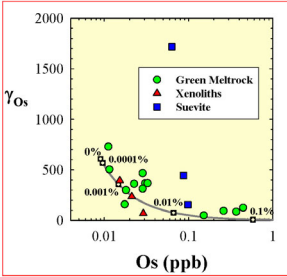


Figure 4: Schematic diagram of a simple crater, after [5]. Shows the different types of impactites and their occurrences.

Figure 5: Meteoritic Component and Mixing Curve



This figures shows that not much meteoritic contamination is needed to significantly change the Os concentration and IC of the target rocks. The end-members used in the mixing model are:

Crust: 0.009 ppb Os, ¹⁸⁷Os/¹⁸⁸Os = 0.9,
Carbonaceous Chondrite: 573 ppb Os, ¹⁸⁷Os/¹⁸⁸Os = 0.127

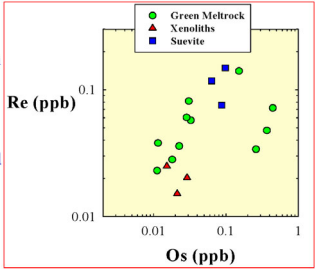
Results

The results suggest that the impact breccias from the depths sampled include only minor Os from the bolide. No samples examined exhibit a pure meteoritic signature. This is consistent with both the low abundances of Os in the rocks, the suprachondritic Re/Os ratios, and the suprachondritic initial ¹⁸⁷Os/¹⁸⁸Os ratios.

The proportion of the impactor is estimated at ~ 0.1%. The model used assumed a ¹⁸⁷Os/¹⁸⁸Os ratio of 0.9 for the target rocks, similar to the isotopic composition of modern seawater [11], the bolide component could constitute as much as 40% of the Os in the meltrocks (less than 0.1% of the total mass of the rock). More likely the isotopic compositions of the target rocks were substantially lower and the percentage of bolide Os is likely <10%. Atom for atom the meteorite contributed more Os, ~ 10%, than the target rocks contained pre-impact. However, as the bolide

Figure 6: Meteoritic component in the different types of impact meltrocks studied.

Suevitic meltrocks had the most significant meteoritic contamination followed by the xenoliths. The green meltrocks had the most variability. These concentrations suggested that while all of the samples contained a meteoritic component, they were variably contaminated.



disintegrated only about 0.1% of the impactites were originally meteoritic material. The lower initial ratio of the suevites and higher Os concentration may suggest either a higher percentage of the bolide component, or a substantially different mix of target lithologies. Effects of subsequent alteration cannot yet be estimated, but the similarity in isotopic systematics between the closely spaced YAX-1 861.4 and 863.51 samples (combined they define a 65 Ma isochron) suggests minimal open-system behavior.

The UNAM samples did come from the YAX-1 borehole. They are also green monomictic impact melt breccias. These samples which are purported to have meteoritic clasts within them mostly have a substantially lower Os isotopic ratio than all of the other types of samples. The ratio is still suprachondritic.

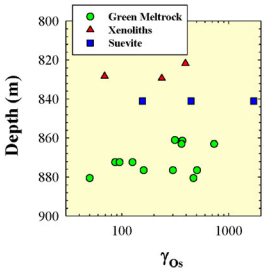
Table 1: Rhenium-Osmium Data for Chicxulub Impact Melt Breccias

Sample	Weight (g)	Re (ppb)	Os (ppb)	¹⁸⁷ Os/ ¹⁸⁸ Os	¹⁸⁷ Re/ ¹⁸⁸ Os	¹⁸⁷ Os/ ¹⁸⁸ Os	¹⁸⁷ Os/ ¹⁸⁸ Os
YAX-1 831.76 (3.33x)	0.36	0.025	0.034	0.034	8.34	0.025	304
YAX-1 828.28 (38.0x)	0.98	0.020	0.029	0.029	3.40	0.023	49
YAX-1 829.36 (34.3x)	0.54	0.015	0.022	0.028	3.56	0.026	237
YAX-1 841.32 (65.4)	0.40	0.149	0.084	0.336	7.47	0.324	156
YAX-1 841.32 (65.5)	0.42	0.117	0.065	2.302	11.4	2.30	1717
YAX-1 841.32 (65.4)	0.31	0.075	0.087	0.040	4.45	0.080	445
YAX-1 861.4 (55.4x)	0.73	0.081	0.087	0.603	13.6	0.992	367
YAX-1 861.4 (55.3)	0.51	0.057	0.036	0.022	8.93	0.592	348
YAX-1 861.4 (55.6)	0.11	nd	0.025	0.536	nd	nd	314
YAX-1 863.51 (67.5)	0.72	0.036	0.026	0.902	8.11	0.84	362
YAX-1 863.51 (67.3)	0.53	0.033	0.012	1.082	10.7	1.05	730
YAX-1 872.4 (UNAM1)	0.89	0.072	0.445	0.268	0.395	0.268	136
YAX-1 872.4 (UNAM1)	0.65	0.048	0.381	0.260	0.634	0.256	86.6
YAX-1 872.4 (UNAM1)	0.72	0.034	0.260	0.249	0.637	0.247	95
YAX-1 876.46 (64.5)	0.12	nd	0.015	0.389	nd	nd	160
YAX-1 876.46 (64.4)	0.54	0.028	0.081	0.516	7.80	0.907	301
YAX-1 876.46 (64.6)	0.46	0.038	0.015	0.847	17.1	0.76	365
YAX-1 880.5 (UNAM2)	0.86	0.040	0.025	0.745	11.0	0.719	488
YAX-1 880.5 (UNAM2)	0.77	0.141	0.156	0.198	4.51	0.190	50

* From the first set of data, contained an error due to the small amount of sample and the blank to spike ratio.

Figure 7: Depth vs. Meteoritic Component

Each of the different types of samples were retrieved from the borehole at approximately the same depths. The meteoritic component does not vary substantially with depth.



Conclusions

- Most rocks were variably contaminated with the impactor.
- None of the samples are purely chondritic.
- Little meteoritic material, 0.1%, is required to significantly change the Os concentration and isotopic composition in the target rocks.
- The bolide only contributed ~ 0.1% of its bulk composition to the impactites.

References: [1] Kyte (1998) Nature 396, 236-239. [2] E. Pierazzo et al. (1998) J. Geophys. Res. 103, 28607-28625. [3] Shukolyukov and Lugmair (1998) Science 282, 927-929. [4] Sharpton et al. (1992) Nature 359, 819-821. [5] Koeberl et al. (1994) Geochim. Cosmochim. Acta 58, 1679-1684. [6] Schuraytz et al. (1994) Geology 22, 868-872. [7] Kring et al. (2003) Lun. and Planet. Sci. XXXIV. [8] Kring, et al. (1991) Lun. and Planet. Sci. XXII, 755-756. [9] Kring and Boynton (1992) Nature 358, 141-144. [10] Zurcher et al. (2003) Lun. and Planet. Sci. XXXIV. [11] Sharma et al. (1997) Geochim. Cosmochim. Acta 61, 5411-5416. [12] Horan et al. (2003) Chem. Geol., in press. [13] Smoliar et al. (1996) Science 271- 1099-1102. [14] Shirey and Walker (1998) Ann. Revs. Earth Planet. Sci. 26, 423-500. [15] Sharpton et al. (1996) G.S.A. Special Paper 307, 55-74. [16] Kring (2000) "The Chicxulub Crater." www.icdp-online.de/html. [17] Alvarez et al. (1980) Science 208, 1095-1108. **Acknowledgements:** This work was partially supported by NASA grant NAG510425 (to R.J.W.). We also wish to acknowledge the International Continental Drilling Program, Universidad Nacional Autonoma de Mexico, and Chicxulub Scientific Drilling Project team.