

## Tectonic interpretation in the Modi Khola valley, central Nepal using zirconium-in-rutile Thermometry

**Todd Reitz GEOL 393** 

Advisors: Aaron Martin, Sarah Penniston-Dorland Phil Piccoli



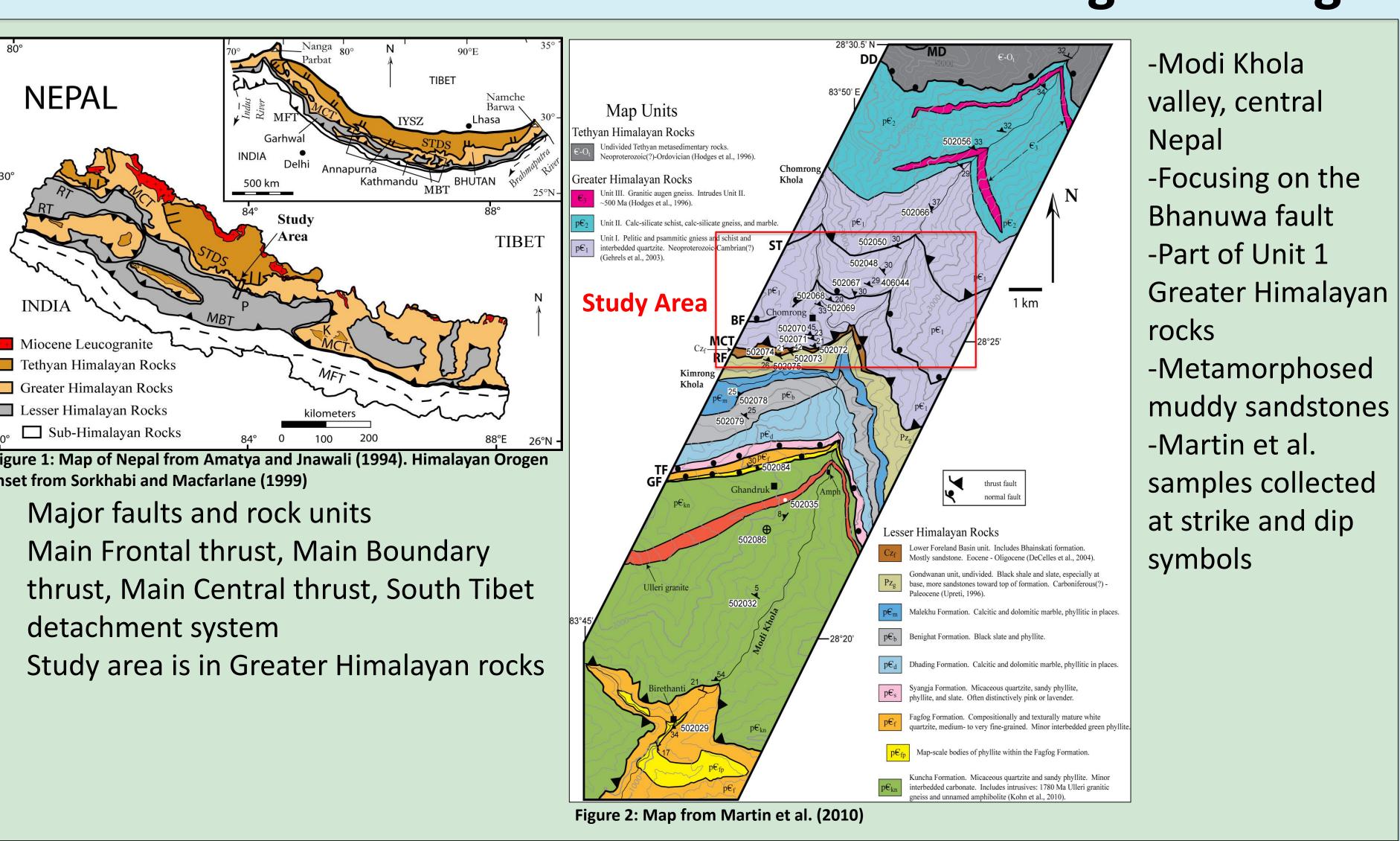
### 1. Introduction

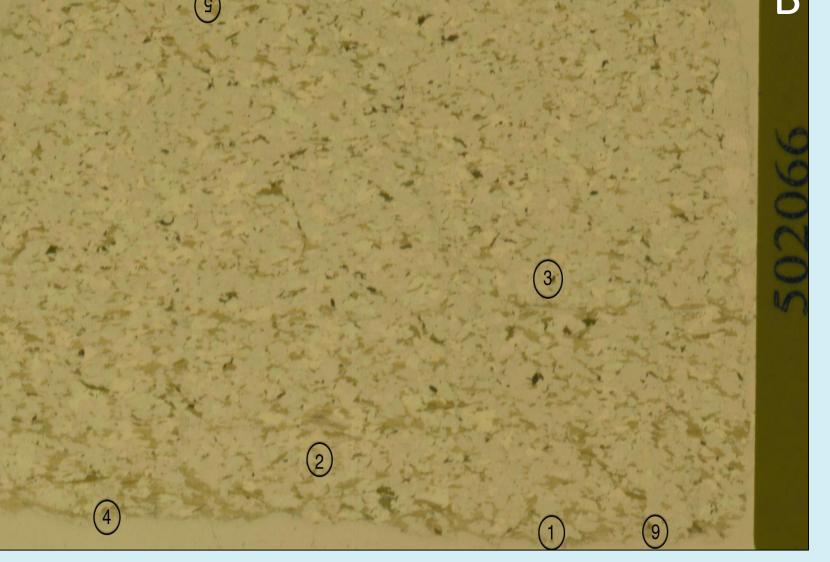
Metamorphism of the Greater and Lesser Himalayan rocks have been studied by Martin et al. (2010) and Corrie and Kohn (2011). Both groups in the Modi Khola valley, central Nepal and used similar thermobarometric methods. The two groups did not produce the same pressure/temperature estimates for the area of study, which has leads to inconsistencies in interpretations of faulting of the region. These discrepancies have implications for the tectonic interpretation of the region, using the zirconium-inrutile thermometer calibrated by Tomkins et al. (2007). Furthermore, the temperature estimates calculated from this thermometer will be used as closure temperatures to compute the relative cooling rates of one side of the fault to the other. A sense of motion across the Bhanuwa fault can then be interpreted.

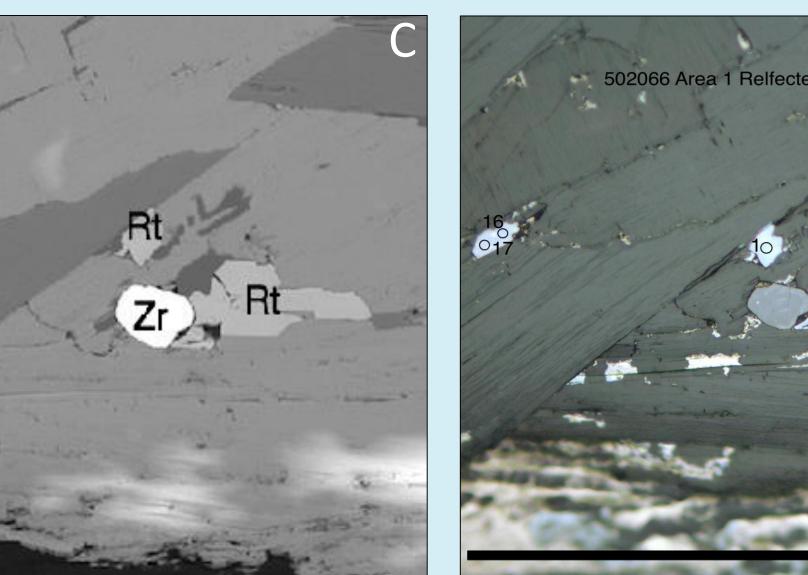
## 4. Hypothesis

Zirconium-in-rutile temperature estimates are consistent within uncertainty with estimates from Martin et al. (2010) and Corrie and Kohn (2011)

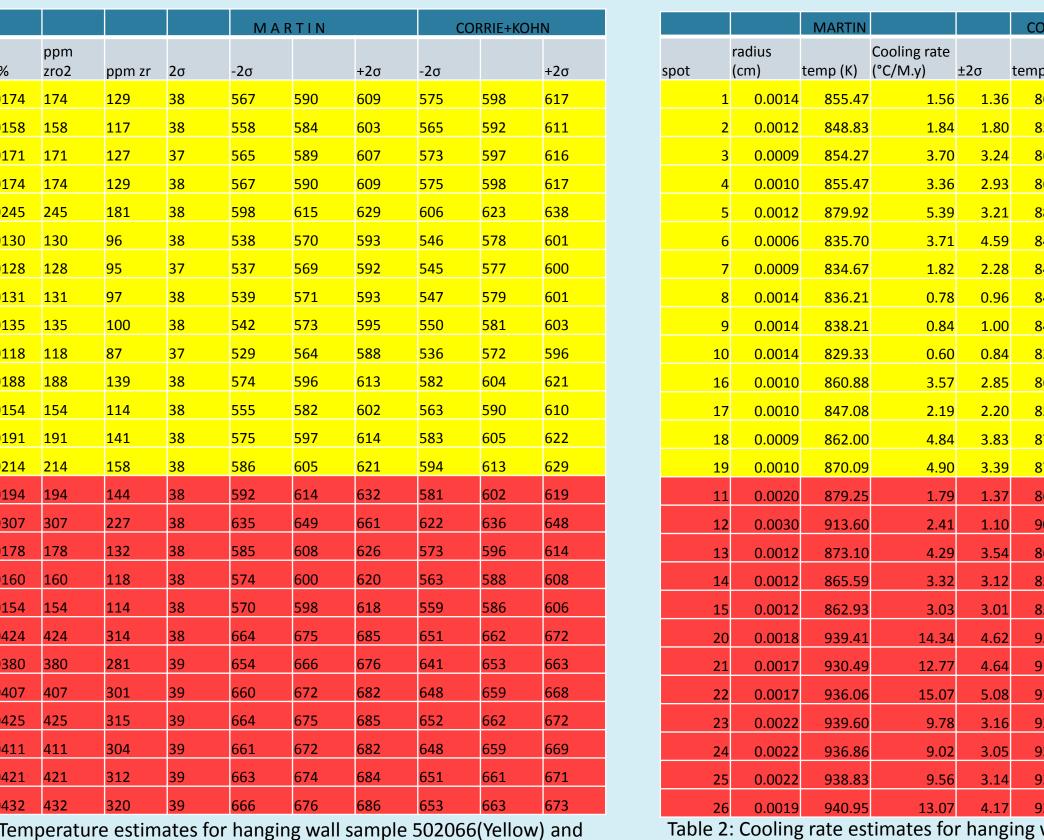
# 2. Geologic Setting







### 8. Results



1. After calculating temperature estimates using the zirconium-in-

rutile thermometer it was noticed that these temperatures are

2. Interpreting the temperature estimates as closure temperatures,

absolute values of the cooling rates are not important, what

they could be plugged into an equation to get cooling rate. The

matters is the relative cooling rates of one side of the fault to the

other side. Assume that the faster cooling rate= faster exhumation

rate and that the transformation from cooling rate to exhumation

normal fault. If the hanging wall exhumed more rapidly, then it was

3. Initial results based on a very small data set are consistent with the

samples in the footwall samples cooling quicker than the hanging

wall samples, suggesting normal fault. This conclusion will be

rate was the same for both the hanging wall and footwall. If the

hanging wall exhumed more slowly than the footwall it was a

lower than other works. We concluded that these rutile grains are

9. Conclusions

a thrust.

too small to record peak temperatures.

explored further next semester.

Journel of metamorphic geology. 25. (2007): 703-713.

temperature estimates from Corrie and Kohn (2011) pressures

### 5. Methods

- A. Analyze thin sections using petrographic microscope. Document the mineral assemblages found and locate rutile grains using petrolographic/optical properties for further analysis using the EPMA. (Photo A above)
- B. Record location of rutile grains on thin section map (Photo B above)
- C. Using the EPMA locate rutile grains (Photo C above)
- E. Plug in the concentrations of zirconium in ppm into the Tomkins et al. (2007) α-quartz calibration of the zirconium-in-rutile thermometer (Formula E)
- (1972) to calculate cooling rates (Formula F)

- D. Choose spots for analysis on grains of rutile to be measured for their zirconium concentrations (Photo D above)
- F. Use estimated closure temperatures calculated and equation 23 from Dodson

## 3. Background

inset from Sorkhabi and Macfarlane (1999)

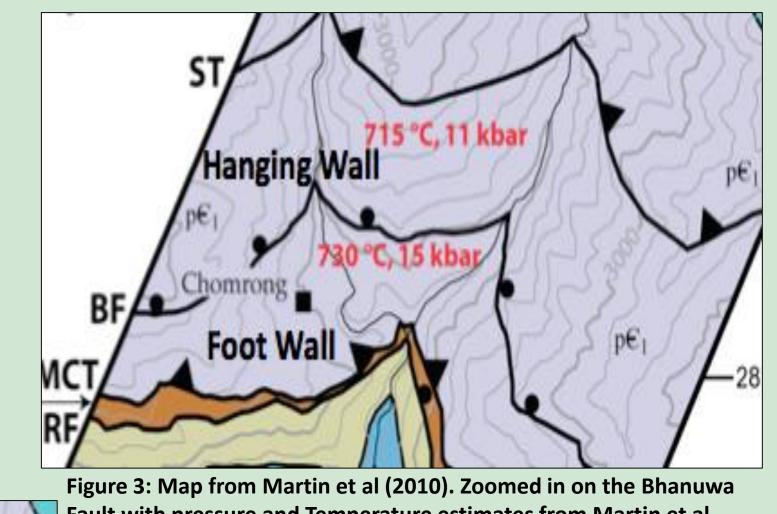
detachment system

Major faults and rock units

Main Frontal thrust, Main Boundary

thrust, Main Central thrust, South Tibet

Martin et al. (2010) overview -thermometer: garnet-biotite -barometer: garnet-muscovitequartz-biotite-plagioclase -4 kbar pressure difference led to interpretation of normal fault



### Corrie and Kohn (2011) overview

-thermometer: garnet-biotite -barometer: garnet-muscovite-quartzbiotite-plagioclase

-85°C temperature difference led to interpretation of thrust fault

### 6. Equations

83.9+0.410*P*  $T(^{\circ}C) = \frac{0.000}{0.1428 - R \ln \phi}$ 

P is pressure in kbar, R is the gas constant (.0083144 kJ/K)

 $\phi$  is the Zr concentration in ppm. Pressures estimated across the fault from Martin et al. (2010) and Corrie and Kohn (2011) will be used.

$$\frac{E}{RT_c} = \ln(-(ART_c^2 D_0)/(Ea^2 \left(\frac{dT}{dt}\right))) \quad \mathsf{F}$$

E is the activation energy

R is the gas constant

T<sub>c</sub> is the closure temperature A is the geometric factor  $(A=e^G)$ 

 $D_0$  is the pre-exponential factor in the Arrhenian expression of diffusion coefficient a is the radius of the grain

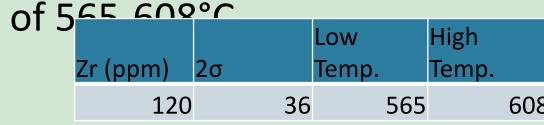
(dT/dt) is the cooling rate.

Using values from Cherniak et al. (2007) for E and D<sub>0</sub>, values given at 170±30 kJ/mol for E and  $9.8E^{-11}$  cm<sup>2</sup>/s for D<sub>0</sub>. It was concluded by Blackburn et al. (2012) that 200kJ/mol is more appropriate value for  $D_0$ .  $T_c$  and a are my estimated values and radius determined from photomicrograph images, the G for sphere given as 4.0066 from Ganguly et al. (2009).

### 7. Uncertainty

Uncertainties on the EPMA can be estimated using standard counting statistical techniques. These uncertainties are estimated from a modified version of the equation  $2\sigma=2((\sqrt{N})/N)*100\%$ . Where N is the number of counts at peak position.

An example given below has an uncertainty of 15% due to counting statistics and a pressure of 12 kbars. This example is measured as 120±36 ppm Zr, which gives a temperature range



Calculated values of cooling rate only show minimal uncertainty, that of the analytical uncertainties from the EPMA. Other uncertainties affect the true absolute value of the cooling rate, but for the purposes of this study in terms of relative cooling rates of hanging wall to footwall those uncertainties cancel out. The uncertainty of grain size, more importantly its radius, needs to be looked at further since it does not cancel out along with the uncertainties from standard counting statistical techniques

References

-Amatya, K.M. and B.M. Jnawali. Geological map of Nepal. Department of Mines and Geology, Government of Nepal, Scale 1:1,000,000. (1994)

-Blackburn, T, N Shimizu, S.A. Bowring, B Schoene, K.H. Mahan. "Zirconium in rutile speedometry: New constraints on lower crustal cooling rates and residence temperatures." Earth and Planetary Science Letters 317-318 (2012): 231-240. -Cherniak, D.J., J Manchester, and E.B. Watson. "Zr and Hf diffusion in rutile." Earth and Planetary Science Letters 261 (2007): 267-279. -Corrie, S.L., and M.J. Kohn. "Metamorphic history of the central Himalaya, Annapurna region, Napal, and implications for tectonic models."

Geological Society of American Bulletin. 123. (2011): 1863-1879. -Dodson, Martin H. "Closure Temperature in Cooling Geochronological and Petrological systems." Contributions to Mineralogy and Petrology, 40 (1972): 259-274. -Ganguly, J and T. Massimilano. "Closure Temperature, Cooling Age and High Temperature Thermochronology." Physics and Chemistry of Earth's Interior; Crust, Mantle, Core. (2009) 89-98. -Martin, Aaron, Jibamitra Ganguly, and Peter DeCelles. "Metamorphism of Greater and Lesser Himalayan rocks exposed in the Modi Khola valley, central Nepal."

Contributions to Mineralogy and Petrology. 159. (2009): 203-223. -Sorkhabi, R.B. and A. Macfarlane." Himalaya and Tibet: mountain roots to mountain tops. In: Macfarlane A, Sorkhabi RB, Quade J (eds) Himalaya and Tibet: mountain roots to mountain tops"Geological Society of America. United States, Geological Society of America (GSA), (1999)Boulder, CO, pp 1–7. -Tomkins, H.S., R Powell, and D.J. Ellis. "The pressure dependence of zirconium-in-rutile thermometer."

**NEPAL** 

